

NPS-53-79-001

# NAVAL POSTGRADUATE SCHOOL

## Monterey, California



A TRANSFER FUNCTION ANALYSIS OF NUMERICAL SCHEMES  
USED TO SIMULATE GEOSTROPHIC ADJUSTMENT

Arthur L. Schoenstadt

Final Report for Period January - October 1978

Approved for Public Release; Distribution Unlimited

Prepared for: Chief of Naval Research  
Arlington, VA 22217

FEDDOCS  
D 208.14/2:  
NPS-53-79-001

NAVAL POSTGRADUATE SCHOOL  
Monterey, California

Rear Admiral Tyler F. Dedman  
Superintendent

Jack R. Borsting  
Provost

The work reported herein was supported by the Foundation Research Program of the Naval Postgraduate School with funds provided by the Chief of Naval Research.

Reproduction of all or part of this report is authorized.

This report was prepared by:

UNCLASSIFIED

SECURITY CLASSIFICATION OF THIS PAGE (When Data Entered)

REPORT DOCUMENTATION PAGE		READ INSTRUCTIONS BEFORE COMPLETING FORM
1. REPORT NUMBER NPS-53-79-001	2. GOVT ACCESSION NO.	3. RECIPIENT'S CATALOG NUMBER
4. TITLE (and Subtitle)  A TRANSFER FUNCTION ANALYSIS OF NUMERICAL SCHEMES USED TO SIMULATE GEOSTROPHIC ADJUSTMENT		5. TYPE OF REPORT & PERIOD COVERED Final Jan - Oct 1978
		6. PERFORMING ORG. REPORT NUMBER
7. AUTHOR(s)  Arthur L. Schoenstadt		8. CONTRACT OR GRANT NUMBER(s)
9. PERFORMING ORGANIZATION NAME AND ADDRESS Naval Postgraduate School Monterey, CA 93940		10. PROGRAM ELEMENT, PROJECT, TASK AREA & WORK UNIT NUMBERS 61152N, RR 000-01-01  N0001478WR80023
11. CONTROLLING OFFICE NAME AND ADDRESS  Naval Postgraduate School Monterey, CA 93940		12. REPORT DATE November 1978
		13. NUMBER OF PAGES
14. MONITORING AGENCY NAME & ADDRESS (if different from Controlling Office)		15. SECURITY CLASS. (of this report)  UNCLASSIFIED
		15a. DECLASSIFICATION/DOWNGRADING SCHEDULE
16. DISTRIBUTION STATEMENT (of this Report)  Approved for Public Release; Distribution Unlimited		
17. DISTRIBUTION STATEMENT (of the abstract entered in Block 20, if different from Report)		
18. SUPPLEMENTARY NOTES		
19. KEY WORDS (Continue on reverse side if necessary and identify by block number)  Finite Difference Methods, Finite Element Methods, Fourier Transform Analysis, Geostrophic Adjustment, Numerical Weather Prediction		
20. ABSTRACT (Continue on reverse side if necessary and identify by block number) Traditional analysis of finite difference and finite element schemes used in numerical weather prediction has concentrated almost exclusively on an analysis of the accuracy with which the scheme simulates the phase speeds involved. In this paper, we use the transfer function formulation common in design of electrical filters to show that these numerical schemes also have important amplitude distortion characteristics, which must be consider- ed in any complete analysis.		



## TABLE OF CONTENTS

I.	INTRODUCTION .....	1
II.	TRANSFER FUNCTIONS .....	2
III.	DISCRETE FILTERS AND TRANSFER FUNCTIONS .....	5
IV.	THE SHALLOW WATER EQUATIONS AS A FILTER .....	7
V.	TRANSFER FUNCTIONS FOR QUASI-DISCRETE FINITE DIFFERENCE AND FINITE ELEMENT METHODS .....	12
VI.	COMPARISON OF THE ARAKAWA SCHEMES .....	17
VII.	CONCLUSIONS .....	20
	ACKNOWLEDGMENTS .....	22
	REFERENCES .....	23
	TABLES .....	24
	LIST OF FIGURES .....	26
	FIGURES .....	27



## I. INTRODUCTION

The main concern when approximating either an ordinary or a partial differential equation by a discrete equation, either finite difference or finite element, is the degree to which the discrete equation solution approximates the solution to the differential equation. The closeness of this approximation has traditionally been considered from both its quantitative aspects, such as relative error (Henrici (1963)), and its qualitative aspects, such as behavior of transients, propagation of fronts, etc. (Haltiner (1971)). Analyses based primarily on consideration of the basically qualitative property of numerical dispersion, where the discretization process causes distortions in the propagation velocities for different frequencies, are especially prevalent (Arakawa and Lamb (1977)).

In this paper, we apply certain ideas derived from the concepts of transfer functions and digital filters in electrical engineering to the study of the qualitative behavior of discretization schemes. The particular model we shall use is based on the linearized, one dimensional shallow water equations, without mean flow, a model which has important application to the study of the process of geostrophic adjustment. (It will be evident that the basic approach, however, is not limited to this model.) The qualitative effects of discretization schemes on this equation have been studied by Winninghoff (1968), Arakawa and Lamb (1977), Schoenstadt (1977), and others. We shall show that use of the transfer function concept leads to important insights into the differences caused by different discretization schemes - differences that are not fully evident from phase propagation considerations only.

## II. TRANSFER FUNCTIONS

It is well known (Stremmler (1977)) that linear, space (or time) invariant systems can be fully described in terms of their so-called transfer function. (Figure 1) That is, if we denote the input to the system as  $y_i(x)$ , and the output as  $y_o(x)$ , then the Fourier transforms of the input and output, denoted respectively as

$$\tilde{y}_i(k) = \int_{-\infty}^{\infty} y_i(x) e^{-ikx} dx \quad , \quad (1)$$

$$\tilde{y}_o(k) = \int_{-\infty}^{\infty} y_o(x) e^{-ikx} dx \quad ,$$

are related by

$$\tilde{y}_o(k) = \tilde{\phi}(k) \tilde{y}_i(k) \quad . \quad (2)$$

Equation (2) is commonly said to represent the system in the transform domain. The equivalent representation in the physical (time or space, as appropriate) domain is

$$y_o(x) = \int_{-\infty}^{\infty} \phi(x - s) y_i(s) ds \quad . \quad (3)$$

In equation (3),  $\phi(x)$  is commonly referred to as the impulse response of the system, and is interpreted as the response of the system to an input delta function at  $x = 0$ , i.e.  $\delta(x)$ .

In general,  $\tilde{\phi}(k)$  is a complex valued function. The physical interpretation of both the magnitude and phase of  $\tilde{\phi}(k)$  is easily arrived at by considering the response of the system to a single sinusoidal input of arbitrary frequency, i.e.

$$y_i(x) = e^{ik_i x} \quad . \quad (4)$$



It is very straightforward to show that

$$y_o(x) = \tilde{\phi}(k_i) e^{ik_i x}.$$

Thus, if we represent  $\tilde{\phi}(k)$  in polar form as

$$\tilde{\phi}(k_i) = |\tilde{\phi}(k_i)| e^{ik_\phi},$$

equation (5) becomes:

$$\begin{aligned} y_o(x) &= |\tilde{\phi}(k_i)| e^{i(k_i x + k_\phi)} \\ &= |\tilde{\phi}(k_i)| e^{ik_i(x + (k_\phi/k_i))}. \end{aligned} \quad (5)$$

Clearly, from (5), we can interpret the magnitude of the transfer function as the factor by which the amplitude of a sinusoid of a given frequency is either amplified or attenuated, and the argument of the transfer function as determining the amount by which the phase of the output sinusoid is shifted relative to the phase of the input sinusoid.

In general, neither  $|\tilde{\phi}(k)|$  nor  $(k_\phi/k)$  is constant, and hence, the different frequencies in a given input are amplified/attenuated and shifted in phase by different amounts. Thus the output signal generally has its shape (graph) altered from the input signal. (Figure 2) This alteration of shape is commonly called distortion, and, based on our discussion, is composed of the effects of two actions - commonly called amplitude distortion, which is due to the deviation of  $|\tilde{\phi}(k)|$  from a constant, and phase (or delay) distortion, due to the deviation of  $k_\phi$  from a linear function of  $k$ . Clearly, to completely understand the effect of a linear system on an input signal, one must know both of these effects.

In electrical engineering, it is common to call any linear, time (or space) invariant device, which can be described by a transfer function as discussed above, a filter. We shall use the terminology hereafter in this paper.

Lastly we note that the energy in an "output" is proportional to

$$\begin{aligned}
 \int_{-\infty}^{\infty} y_o^2(x) \, dx &= \frac{1}{2\pi} \int_{-\infty}^{\infty} |\tilde{y}_o(k)|^2 \, dk \quad , \\
 &= \frac{1}{2\pi} \int_{-\infty}^{\infty} |\tilde{\phi}(k) \tilde{y}_i(k)|^2 \, dk \\
 &\leq \frac{1}{2\pi} \left\{ \int_{-\infty}^{\infty} |\tilde{\phi}(k)|^2 \, dk \right\} \left\{ \int_{-\infty}^{\infty} |y_i(k)|^2 \, dk \right\} \quad , \\
 &= \left\{ \int_{-\infty}^{\infty} |\tilde{\phi}(k)|^2 \, dk \right\} \int_{-\infty}^{\infty} y_i^2(x) \, dx
 \end{aligned} \tag{6}$$

and thus one can also use the transfer function to describe the relation between the energies contained in the "input" and "output" to a linear, invariant system. (Note the quantity  $|\tilde{y}(k)|^2$  is commonly referred to as the spectral density of the signal  $y(x)$ .)

### III. DISCRETE FILTERS AND TRANSFER FUNCTIONS

A great deal of research is presently directed toward the study of so-called discrete, or digital filters, which arise when the continuous processes, such as differentiation and integration, in an invariant linear filter are replaced by analogous approximations based on sampled, or discrete data. Thus, for example,

$$y(x + \Delta x) = y(x) + f(x) (\Delta x) \quad , \quad (7)$$

is a digital filter approximating, in some sense (specifically by using an Euler forward predictor), the continuous integrator:

$$\frac{dy}{dx} = f(x) \quad . \quad (8)$$

The transfer functions of such filters can be easily determined using Fourier transforms, and the observation that

$$\int_{-\infty}^{\infty} y(x + \Delta x) e^{-ikx} dx = e^{ik(\Delta x)} \tilde{y}(k) \quad . \quad (9)$$

Thus, the transfer function of the process given by (7) is (where  $f(x)$  is considered the "input" and  $y(x)$  the "output"), after some algebra,

$$\tilde{\phi}(k) = \frac{(\Delta x) e^{ik(\Delta x)/2}}{2i \sin(k(\Delta x)/2)} \quad .$$

(Note the transfer function of the continuous process is  $\tilde{\phi}^c(k) = \frac{1}{ik}$  .)

There is, however, one significant difference between the transfer functions of a continuous filter and its digital (discrete) analog. This difference is commonly referred to with the terms aliasing or folding, and the basic work on it is commonly attributed to Nyquist (Clark (1977)). In its simplest interpretation, Nyquist's work implies that a sampling device, sampling at intervals of  $(\Delta x)$ , is incapable of resolving waves of frequency greater than  $(\pi/(\Delta x))$ , and, if any energy is actually present in a sampled continuous signal at these higher frequencies, it will be erroneously resolved (aliased) into a frequency lower than  $(\pi/(\Delta x))$ .

Thus, while a continuous filter can, in principle, react to any input frequency, and hence has a transfer function defined for all  $k$ , that of a digital filter is considered only for (complex) frequencies between  $\pm(\pi/(\Delta x))$ , the so-called Nyquist cut-off or Nyquist limit.

Lastly, we note that if a digital filter has a significant amplitude response (i.e. if  $|\hat{\phi}(k)|$  is not close to zero) near the Nyquist limit, then the output of that system may have a significant portion of its energy in these high frequency oscillations. These oscillations often appear to the observer as noise, or "noise-like", and hence, in discussing digital filter design, many books and articles recommend controlling the transfer function response near this limit. (This is sometimes referred to as "windowing".)

#### IV. THE SHALLOW WATER EQUATIONS AS A FILTER

The basic equations we propose to study, using the ideas discussed above, are the linearized, one-dimensional, shallow water equations in an infinite region, with no mean flow:

$$\left. \begin{aligned} \frac{\partial u}{\partial t} - fv + g \frac{\partial h}{\partial x} &= 0 \\ \frac{\partial v}{\partial t} + fu &= 0 \\ \frac{\partial h}{\partial t} + H \frac{\partial u}{\partial x} &= 0 , \end{aligned} \right\} \quad (10)$$

where  $u$  denotes the perturbation velocity in the  $x$  direction,  $v$  the perturbation velocity normal to the  $x$  direction,  $H$  and  $h$  the mean and perturbed heights of the free surface, respectively, and  $g > 0$  and  $f > 0$  the gravitational and Coriolis parameters, respectively. This model is especially important in the study of the meteorological process called geostrophic adjustment, which has been studied in some detail from several approaches by Rossby (1937), Cahn (1945), Winninghoff (1968), Blumen (1972), Arakawa and Lamb (1977) and Schoenstadt (1977). This process of geostrophic adjustment is important because it is the primary mechanism by which an atmosphere, modeled by the so-called meteorological primitive equations, reacts to errors (either numerical or due to errors in observational data) in the initial data. The dispersive wave nature of (10) is the primary mechanism by which these errors (commonly referred to as "imbalances") eventually spread out over the domain of interest, until the system reaches a so-called "balanced" or "geostrophically adjusted" condition. Forecasts at times before this balance is reached are generally inaccurate. Thus a primary consideration in numerically modeling (10) is to insure that the numerical balance mechanism both fairly accurately models the physical balancing pro-

cedure, and that it proceeds relatively quickly.

As discussed in Schoenstadt (1977), equations (10) can be easily solved by using spatial Fourier transform. If we continue to denote Fourier transforms by an overhead tilde, e.g.:

$$\tilde{u}(k,t) = \int_{-\infty}^{\infty} u(x,t) e^{-ikx} dx, \quad (11)$$

etc. Then (10) reduces directly to:

$$\left. \begin{aligned} \frac{d\tilde{u}}{dt} &= f\tilde{v} - ikg\tilde{h}, \\ \frac{d\tilde{v}}{dt} &= -f\tilde{u}, \\ \frac{d\tilde{h}}{dt} &= -ikH\tilde{u}, \end{aligned} \right\} \quad (12)$$

together with initial conditions:

$$\tilde{u}_i = \tilde{u}(k,0) = \int_{-\infty}^{\infty} u(x,0) e^{-ikx} dx, \quad (13)$$

etc. Then since (12) represents simply a set of coupled, constant coefficient (in the sense of  $t$ ) ordinary differential equations, their solution can be easily shown, by the usual eigenvalue/eigenvector approach, to be:

$$\left. \begin{aligned} \tilde{u}(k,t) &= \tilde{u}_i \cos vt + \frac{f\tilde{v}_i}{v} \sin vt - \frac{ikg\tilde{h}_i}{v} \sin vt, \\ \tilde{v}(k,t) &= -\frac{f}{v} \tilde{u}_i \sin vt + \left\{ \frac{k^2 gH}{v^2} + \frac{f^2}{v^2} \cos vt \right\} \tilde{v}_i + \frac{ikgf}{v^2} \{1 - \cos vt\} \tilde{h}_i, \\ \tilde{h}(k,t) &= -\frac{ikH}{v} \tilde{u}_i \sin vt - \frac{ikHf}{v^2} \{1 - \cos vt\} \tilde{v}_i + \left\{ \frac{f^2}{v^2} + \frac{k^2 gH}{v^2} \cos vt \right\} \tilde{h}_i \end{aligned} \right\} \quad (14)$$

where:

$$v^2 = f^2 + k^2 gH = f^2(1 + \lambda^2 k^2), \quad \lambda = \sqrt{gH}/f. \quad (15)$$

We now consider (14) in view of our discussion of filters and transfer functions. Observe that if we denote

$$\underline{y}_i = \begin{bmatrix} u_i \\ v_i \\ h_i \end{bmatrix} \quad \text{and} \quad \underline{y}_0 = \begin{bmatrix} u \\ v \\ h \end{bmatrix} \quad (16)$$

respectively, as the "input" and "output" to the "system" described by (10), then we can express (14) as

$$\underline{\tilde{y}}_o = \underline{\tilde{\phi}(k,t)} \underline{\tilde{y}}_i \quad (17)$$

where,

$$\underline{\tilde{\phi}(k,t)} = \begin{bmatrix} \cos \nu t & \frac{f_o}{\nu} \sin \nu t & \frac{ikg_o}{\nu} \sin \nu t \\ -\frac{f}{\nu} \sin \nu t & \left\{ \frac{k^2 gH}{\nu^2} + \frac{f^2}{\nu^2} \cos \nu t \right\} & \frac{ikgf}{\nu^2} \{1 - \cos \nu t\} \\ -\frac{ikH}{\nu} \sin \nu t & -\frac{ikHf}{\nu^2} \{1 - \cos \nu t\} & \left\{ \frac{f^2}{\nu^2} + \frac{k^2 gH}{\nu^2} \cos \nu t \right\} \end{bmatrix} \quad (18)$$

Immediate comparison of (17)-(18) with (2) is not possible since  $\underline{\tilde{\phi}(k,t)}$  depends not only on the transform variable ("k"), but also on one of the physical variables ("t"). However, since  $\sin \nu t$  and  $\cos \nu t$  can be expressed in terms of  $e^{\pm i\nu t}$ , and since  $k$  is real, it is easily seen that the sinusoidal terms in (18) express only the phase relationships (which clearly change with time), while the other, time independent, coefficients reflect the amplitude effects of the "filter". In fact, it is fairly easy to show (Schoenstadt (1977)) that the sinusoidal terms in (18) produce dispersive, transient waves, propagating throughout the region. (These waves are clearly dispersive since  $(\nu(k)/k)$  is not constant.) The time independent coefficients in (18) still function as amplitude distortion terms in the sense that they redistribute the energy that is in the input (initial condition) waves, in a manner that varies with wave number, into the output (solution) waves.

A close examination of (18) shows that, in fact, the "amplitude distortion" in this "system" is really governed in each term by precisely one of



the three factors,

$$\frac{1}{v} , \quad \frac{k}{v} \quad \text{or} \quad \frac{k}{v^2} ,$$

or the square of one of these. The amplitude response curves for these three expressions are shown in Figure 3. It is clear that the terms with the coefficient  $(1/v)$  have their low frequencies least affected, the terms with coefficient  $(k/v)$  their high frequencies least affected, and those with coefficient  $(k/v^2)$  their frequencies in some middle range least effected. Such responses are often referred, respectively, to as low pass, high pass, and band pass filters.

Since the "system" described by (10) involves both spatial and temporal variation, it is well known that not only the phase velocity  $(v(k)/k)$  , but also the group velocity  $(\frac{dv}{dk})$  is important in describing the behavior, since it is the latter which governs the rate at which energy actually propagates. Thus, in Figure 4, we display both of these velocities for the system (10).

We note that additional insight into some of the qualitative behavior of the solution to (10) can be obtained by decomposing (18) into its transient and steady state components, i.e., writing

$$\tilde{y}_o = \tilde{\phi}_S \tilde{y}_i + \tilde{\phi}_T \tilde{y}_i , \quad (19)$$

where  $\tilde{\phi}_S$  and  $\tilde{\phi}_T$  can be obtained by inspection, since only time dependent terms contribute to the transients, however, we shall not pursue this analysis here.

As a final comment, note that the total energy of the system represented by (10) is given by:

$$u^2 + v^2 + \frac{g}{H} h^2 .$$

Thus the spectral density of the steady state field can be shown to be

$$|u_S|^2 + |v_S|^2 + \frac{g}{H} |h_S|^2 = \frac{gHk^2}{v^2} |v_i|^2 + \frac{gf^2}{Hv} |h_i|^2 . \quad (20)$$



Since the term  $(k/v)$  acts as a high-pass filter, and the term  $(f/v)$  as a low pass-filter, we see that the energy in the steady state field is determined by a combination depending primarily on the high frequency energy in the initial  $v$  field, and the low frequency energy in the initial  $h$  field.

We shall not proceed any further with this analysis, since our main interest is to analyze how discrete schemes, and the corresponding "filters" they produce, approximate the continuous (differential) model. However, there are certainly other insights into the general behavior of the physical model still possible from this filter point of view.

## V. TRANSFER FUNCTIONS FOR QUASI-DISCRETE FINITE DIFFERENCE AND FINITE ELEMENT METHODS

In this section, we extend the analysis of our continuous model to the quasi-discrete case, where continuous spatial operations are replaced by operations only at sample points, while the continuous time operations are retained. This approach follows that in Winninghoff (1968) and Arakawa and Lamb (1977). However, we extend our analysis to include a more general case of quasi-discrete schemes than they considered. Specifically, we define an invariant linear discretization operator about the point  $x$  as any operator having the form

$$L[f(x)] = \sum_{n=-n_1}^{n_2} c_n f(x + n(\Delta x)) , \quad (21)$$

where the  $c_n$  denote real constants.

The discretization schemes we shall investigate are specifically those satisfying the following conditions:

- (a) All discretization operators are linear and time and space invariant.
- (b) All discretization operators operate only on symmetric sets of sample points, i.e. in (21),  $n_1 = n_2$ .
- (c) The discretization operators are balanced, i.e. in (21)  $|c_n| = |c_{-n}|$ .
- (d) The operators are applied to (8)-(10) consistently, i.e. if one of the derivative terms on the right hand side of (8)-(10) is replaced by a certain difference, then all derivative terms on the right hand side of (8)-(10) are replaced by the same difference operator.
- (e) The discretization must be exact, i.e. produce exactly the same result as the analogous continuous operation, on all at least arbitrary linear functions, i.e.  $f(x) = ax + b$ ,  $a, b$ , arbitrary.

With these restrictions, it is fairly simple to show that the quasi-discrete schemes we wish to consider can all be reduced to the general form:

$$\left. \begin{aligned} \frac{\partial}{\partial t} \{L_{\alpha}[u(x,t)]\} &= f L_{\beta}[v(x,t)] - g L_{\sigma}[h(x,t)] \\ \frac{\partial}{\partial t} \{L_{\alpha}[v(x,t)]\} &= -f L_{\beta}[u(x,t)] \\ \frac{\partial}{\partial t} \{L_{\alpha}[h(x,t)]\} &= -H L_{\sigma}[u(x,t)] \end{aligned} \right\} \quad (22)$$

where  $L_{\alpha}[\ ]$ ,  $L_{\beta}[\ ]$ , and  $L_{\sigma}[\ ]$ , are invariant linear discretization operators having the following general form:

$$\left. \begin{aligned} L_{\alpha}[f(x)] &= \alpha_0 f(x) + \sum_{n=1}^{n_{\alpha}} \alpha_n [f(x + n\Delta x) + f(x - n\Delta x)] \\ L_{\beta}[f(x)] &= \beta_0 f(x) + \sum_{n=1}^{n_{\beta}} \beta_n [f(x + n\Delta x) + f(x - n\Delta x)] \\ L_{\sigma}[f(x)] &= \frac{1}{\Delta x} \sum_{n=1}^{n_{\sigma}} \sigma_n [f(x + n\Delta x) - f(x - n\Delta x)] \end{aligned} \right\} \quad (23)$$

The various values of  $\alpha_n$ ,  $\beta_n$ , and  $\sigma_n$  will be referred to as the filter weights. Thus, for example, the finite difference mesh which Arakawa and Lamb (1977) call Scheme C (Figure 5), and which is given in finite difference form as

$$\begin{aligned} \frac{\partial u}{\partial t} &= f \frac{v(x+d/2,t) + v(x-d/2,t)}{2} - g \frac{h(x+d/2,t) - h(x-d/2,t)}{d} \\ \frac{\partial v}{\partial t} &= -f \frac{u(x+d/2,t) + u(x-d/2,t)}{2} \\ \frac{\partial h}{\partial t} &= -H \frac{u(x+d/2,t) - u(x-d/2,t)}{d} \end{aligned}$$

corresponds to the following filter weights:

$\alpha_0 = 1$ ,  $\beta_1 = \frac{1}{2}$ ,  $\sigma_1 = \frac{1}{2}$ , all others zero. (Note that in this formulation  $\Delta x = d/2$ .)

Determination of the filter properties of the system represented by (22) requires taking the Fourier transform (in  $x$ ) of those equations. However, the

shift property of these transforms (equation (9)) simplifies the terms involving the discretization operators, since, using (9)

$$\{L[f(x)]\} = \left\{ \sum_{n=-n_1}^{n_2} c_n e^{ikn(\Delta x)} \right\} \tilde{f}(k) . \quad (24)$$

Thus, the Fourier transform of (22) is

$$\left. \begin{aligned} \alpha(k) \frac{d\tilde{u}}{dt} &= f \beta(k) \tilde{v} - i g \sigma(k) \tilde{h} , \\ \alpha(k) \frac{d\tilde{v}}{dt} &= -f \beta(k) \tilde{u} , \\ \alpha(k) \frac{d\tilde{h}}{dt} &= -i H \sigma(k) \tilde{u} , \end{aligned} \right\} \quad (25)$$

where

$$\alpha(k) = \alpha_0 + 2 \sum_{n=1}^n \alpha_n \cos(nk\Delta x) , \quad (26)$$

$$\beta(k) = \beta_0 + 2 \sum_{n=1}^n \beta_n \cos(nk\Delta x) , \quad (27)$$

$$\sigma(k) = \frac{2}{\Delta x} \sum_{n=1}^n \sigma_n \sin(nk\Delta x) . \quad (28)$$

Equations (25) can be solved, using the same eigenvalue/eigenvector approach as was used for (12), to yield:

$$\left. \begin{aligned} \tilde{u} &= \tilde{u}_0 \cos \hat{v}t + \frac{f\beta\tilde{v}_0}{v} \sin \hat{v}t - \frac{i\sigma g\tilde{h}_0}{v} \sin \hat{v}t \\ \tilde{v} &= -\frac{f\beta\tilde{u}_0}{v} \sin \hat{v}t + \left\{ \frac{\sigma^2 gH}{v^2} + \frac{f^2\beta^2}{v^2} \cos \hat{v}t \right\} \tilde{v}_0 + \frac{i\sigma g f \beta}{v^2} \{1 - \cos \hat{v}t\} \tilde{h}_0 \\ \tilde{h} &= -\frac{i\sigma H \tilde{u}_0}{v} \sin \hat{v}t - \frac{i\sigma H f \beta}{v^2} \{1 - \cos \hat{v}t\} \tilde{v}_0 + \left\{ \frac{f^2\beta^2}{v^2} + \frac{\sigma^2 gH}{v^2} \cos \hat{v}t \right\} \tilde{h}_0 \end{aligned} \right\} \quad (29)$$

where now

$$\hat{v}^2 = \frac{f^2\beta^2(k) + gH\sigma^2(k)}{\alpha^2(k)} ,$$

or

$$\begin{aligned}\hat{v} &= \frac{f}{\alpha(k)} \sqrt{\beta^2(k) + \lambda^2 \sigma^2(k)}, \\ &= \frac{v(k)}{\alpha(k)},\end{aligned}\tag{30}$$

where

$$v(k) = f \sqrt{\beta^2(k) + \lambda^2 \sigma^2(k)}.\tag{31}$$

Comparison of (29)-(31) with (14)-(15) show that the main impacts of the quasi-discretization described by (22) are:

(1) The filter coefficients

$$\frac{1}{v}, \frac{k}{v}, \frac{k}{v^2}$$

and the squares of the first two, are replaced, respectively,

$$\frac{\beta}{v}, \frac{\sigma}{v}, \frac{\beta\sigma}{v^2}$$

and the squares of the first two of these expressions.

(2) The phase velocity

$$\frac{v}{k} = \frac{f}{k} \sqrt{1 + \lambda^2 k^2}$$

is replaced by

$$\frac{f}{k\alpha(k)} \sqrt{\beta^2 + \lambda^2 \sigma^2}.$$

Since the filter coefficients discussed above determine how the energy in the initial disturbance is partitioned into the transient and steady state fields, then the degree to which the discrete filter coefficients approximate the continuous ones can be viewed as a measure of the degree to which the discretization accurately portrays the continuous model energy distribution. Any difference between a continuous (differential) coefficient, and the corresponding discrete coefficient, should be considered as introducing a distortion between the continuous and discrete solutions.

Similarly, any differences between the continuous phase and group velocities and the discrete ones will result in distortion due to spreading of the waves. This effect will be especially noticeable if the distortion is pronounced at (spatial) frequencies that carry significant energy, i.e. for which filter coefficients are "large".

Actually, as noted before, the phase velocity is less important than the group velocity,

$$\frac{dv}{dk}$$

in the continuous case, and

$$\frac{d\hat{v}}{dk}$$

in the discrete case, since this quantity measures the velocity at which the energy in the different frequencies propagates.

## VI. COMPARISON OF THE ARAKAWA SCHEMES

The specific discretization schemes which we shall analyze using (22)-(30) are based on the arrangements of sample points described in Arakawa and Lamb (1977) as Schemes A-D. The arrangements of these points in one dimension is shown in Figure 5. We shall consider both finite difference and finite element formulations of the discrete problem, whereas Arakawa and Lamb considered only second order finite differences. The finite difference formulations are derived for both second and fourth order approximations to the spatial derivative terms, while the finite element formulations are derived (for Schemes A-C only) using piecewise linear elements (the so-called "hat" function basis). The filter weights derived for each of these methods are shown in Tables 1-3. Note that since each scheme "samples" the initial disturbances at a distance of  $d$ , then the shortest wave that can be resolved in each case is of frequency  $(\pi/d)$ .

In Figures 6-8, the amplitude coefficients of the discretized approximate "filters" are compared to the "filters" in the differential model, for the values of the physical parameters given in Arakawa and Lamb (1977). The following qualitative generalizations are clear:

- (1) In each of the methods (second order difference, fourth order difference, or finite elements), the arrangement of points specified by Scheme A appears to produce the greatest distortion, and, furthermore, this distortion is most pronounced at high frequencies (short waves).
- (2) In each of the methods, Schemes B, C and D very accurately approximate the transfer function for the "high pass" filter.
- (3) In each of the methods, Scheme C understates the amount of energy distributed into the short waves by both the low-pass and band pass filters, while Scheme B overstates this.



(4) In all cases, the finite element formulation appears more "accurate" than fourth order differences, which in turn is more "accurate" than the second order differences.

(5) The (group) velocity distortion is especially serious in Schemes A and D, since for short waves there is actually a reversal in the direction of propagation.

Note that even though Scheme B seems to have the least "distortion" in each case, there is still significant phase and group velocity distortion near the cut-off. Thus, we should not conclude immediately that Scheme B will produce the least "noisy" solutions. This is because the apparent noise in the solution is a combination of both amplitude and phase behavior, and any phase distortion may lead to perceived "noise" in the solution if (as is the case in Scheme B) significant energy is retained (due to the amplitude distortion effects) in the short waves. In fact, it may be preferable to tolerate more amplitude distortion, if it acts to reduce the amount of energy in the short waves, which are generally the slowest propagating, and which produce the appearance of noise in the solutions.

As a final check on our analysis, we compared the height fields for both the continuous and all the discrete models, using the same methodology as in Arakawa and Lamb (1977). An initial disturbance given by:

$$u_i(x,0) = 0$$

$$h_i(x,0) = 0$$

$$v_i(x,0) = \begin{cases} v_0 & , -a < x < a \\ 0 & , \text{otherwise} \end{cases}$$

was considered, with the physical parameters having the following values

$$g = 10 \text{ m sec}^{-2}$$

$$H = 10^3 \text{ m}$$

$$f = 10^{-4} \text{ sec}^{-1}$$

$$a = d = \lambda/2.$$



The transform of the initial  $v$  field

$$\tilde{v}_i = \frac{2V_0 \sin(ka)}{k}$$

was then band limited to the region  $-\frac{\pi}{d} < k < \frac{\pi}{d}$ , and the exact solution

for  $h$  was computed as the inverse transform of

$$-\frac{ikHf}{v^2} \{1 - \cos vt\} \left\{ \frac{2V_0 \sin(ka)}{k} \right\}$$

in the continuous (differential) case, and the inverse transform of

$$-\frac{i\sigma\beta fH}{v^2} \left\{ 1 - \cos\left(\frac{v}{\alpha} t\right) \right\} \left\{ \frac{2V_0 \sin(ka)}{k} \right\}$$

in each of the discrete cases. In each instance, the inversion was accomplished using Simpson's rule on the interval  $(0, \frac{\pi}{d})$ , with 600 subintervals.

The different  $h$  fields are shown at Figure 9. Note that these agree with our predictions, based on the filter analysis, in that:

(1) In the second order finite difference methods, Scheme A displays a great deal of numerical noise, as predicted due to the great amplitude distortion of its filter near  $(\pi/d)$ . Schemes B and D show somewhat less noise than Scheme A, but still more than Scheme C, this due to the manner in which Scheme C controls the amplitude near  $(\pi/d)$  in both the  $(\sigma/v)$  and  $(\beta\sigma/v^2)$  terms.

(2) Similar comments hold for the fourth order finite difference methods, though, in general, due to better phase propagation, the spreading of the energy in the short waves is not so pronounced, and hence the noise in Schemes B and D is reduced.

(3) In the finite element methods, Scheme A still remains quite noisy, due to its amplitude distortion near  $(\pi/d)$ . Schemes B and C are virtually identical, since Scheme B seems to compensate for slightly more energy in the short waves by maintaining better phase relationship.

## VII CONCLUSIONS

In this paper we have adapted concepts that are commonly associated with the design of filters in electrical engineering to the analysis of discretized representations of partial differential equations, and specifically to the one dimensional shallow water equations. We have shown that these concepts can provide valuable insights into why certain solution schemes provide "better" solutions than other, and a useful tool for comparing both qualitative and quantitative aspects of different schemes.

We have clearly demonstrated that, in contrast to the analysis used in earlier studies, a complete analysis of the suitability of a certain discrete scheme to model a physical process, in our case the process of geostrophic adjustment, depends not only on providing reasonably accurate duplication of the phase propagation (delay) characteristics of the modeled system, but also on producing reasonably accurate duplication of the amplitude response characteristics. Furthermore, because of the phase distortion near the Nyquist cut-off of discrete schemes, it is often desirable to control the amplitudes of short waves.

In the process of geostrophic adjustment studies, we have concluded that schemes that use unstaggered grid points are generally poor, irrespective of the particular finite difference/finite element method used, due to some inherent properties of the model. The difficulties which have been reported with these methods in the literature, and especially the problem of "noise" in finite element models, appears directly attributable to the amplitude response near the Nyquist cut-off methods based on an unstaggered arrangement of points. This is certainly related to the tendency of schemes based on unstaggered grid points to produce solution separation, and appears to arise from the fact that the coupling of the fields in the basic differential

equation simply makes specifying all the quantities at every grid point an overspecification of the problem.

Lastly, we conclude, that, based on the one dimensional evidence only, the method identified by Arakawa and Lamb (1977) as Scheme C has been quite popular, not only due to its good phase propagation characteristics, but also due to its tendency, in either finite difference or finite element formulations, to "window" out much of the high frequency noise in the discretized model.

In summary, the concepts we have applied yield extremely valuable insights into the qualitative behavior of discrete methods for approximating partial differential equations. In many ways, we have only started to bring much of the emerging understanding of digital filters to bear on this problem. A significant, and interesting outgrowth will be the consideration of two dimensional problems.

## ACKNOWLEDGMENTS

The author is indebted to Professor R. T. Williams for several helpful discussions, and for deriving the filter weights for the finite element schemes presented in Table 3. Computations were performed by the W. R. Church Computer Center, NPS.

## REFERENCES

- Arakawa, A. and Lamb, V. R., "Computational Design of the Basic Dynamical Processes of the UCLA General Circulation Model", Methods of Computational Physics, Vol. 17, Academic Press, 1977.
- Blumen, W., "Geostrophic Adjustment," *Reviews of Geophysics and Space Physics*, Vol. 10, No. 2, May 1972, pp. 485-528.
- Cahn, A., "An Investigation of the Free Oscillations of a Simple Current System," *Journal of Meteorology*, Vol. 2, No. 2, 1945, pp. 113-119.
- Clark, A. P., Advanced Data Transmission Systems, Halsted Press (Wiley), 1977.
- Haltiner, G. J., Numerical Weather Prediction, John Wiley and Sons, Inc., 1971.
- Henrici, P., Error Propagation for Finite Difference Methods, John Wiley and Sons, Inc., 1963.
- Rossby, C. G., "On the Mutual Adjustment of Pressure and Velocity Distributions in Certain Simple Current Systems," *J. Mar. Res.*, 1, 1937-1938, pp. 15-28, 239-263.
- Schoenstadt, A., "The Effect of Spatial Discretization on the Steady-State and Transient Behavior of a Dispersive Wave Equation," *J. Comp. Phys.*, Vol. 23, April 1977, pp. 364-379.
- Stremmler, F. G., Introduction to Communication Systems, Addison-Wesley, 1977.
- Winninghoff, F., "On the Adjustment Toward a Geostrophic Balance in a Simple Primitive Equation Model with Application to the Problem on Initialization and Objective Analysis," *Doctoral Dissertation*, UCLA, 1968.

TABLE 1

FILTER WEIGHTS FOR SECOND ORDER FINITE DIFFERENCES

 $(\Delta x = d/2)$ 

SCHEME	DENOTED	$\alpha_0$	$\beta_0$	$\beta_1$	$\sigma_1$	$\sigma_2$
A	A2	1	1	0	0	$\frac{1}{4}$
B	B2	1	1	0	$\frac{1}{2}$	0
C	C2	1	0	$\frac{1}{2}$	$\frac{1}{2}$	0
D	D2	1	0	$\frac{1}{2}$	0	$\frac{1}{4}$

TABLE 2

FILTER WEIGHTS FOR FOURTH ORDER FINITE DIFFERENCES

 $(\Delta x = d/2)$ 

SCHEME	DENOTED	$\alpha_0$	$\beta_0$	$\beta_1$	$\sigma_1$	$\sigma_2$	$\sigma_3$	$\sigma_4$
A	A4	1	1	0	0	$\frac{8}{24}$	0	$-\frac{1}{24}$
B	B4	1	1	0	$\frac{27}{48}$	0	$-\frac{1}{48}$	0
C	C4	1	0	$\frac{1}{2}$	$\frac{27}{48}$	0	$-\frac{1}{48}$	0
D	D4	1	0	$\frac{1}{2}$	0	$\frac{8}{24}$	0	$-\frac{1}{24}$

TABLE 3

## FILTER WEIGHTS FOR PIECEWISE LINEAR FINITE ELEMENTS

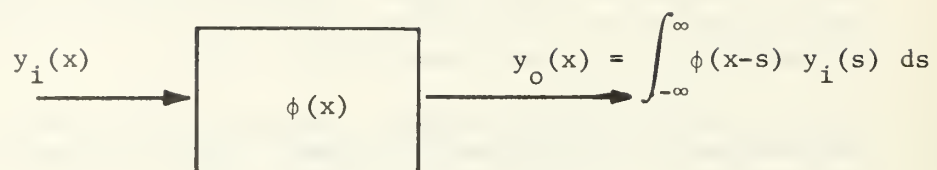
 $(\Delta x = d/2)$ 

SCHEME	DENOTED	$\alpha_0$	$\alpha_1$	$\alpha_2$	$\beta_0$	$\beta_1$	$\beta_2$	$\beta_3$	$\sigma_1$	$\sigma_2$	$\sigma_3$
A	AFE	$\frac{4}{6}$	0	$\frac{1}{6}$	$\frac{4}{6}$	0	$\frac{1}{6}$	0	0	$\frac{1}{4}$	0
B	BFE	$\frac{4}{6}$	0	$\frac{1}{6}$	$\frac{4}{6}$	0	$\frac{1}{6}$	0	$\frac{5}{16}$	0	$\frac{1}{16}$
C	CFE	$\frac{4}{6}$	0	$\frac{1}{6}$	0	$\frac{23}{48}$	0	$\frac{1}{48}$	$\frac{5}{16}$	0	$\frac{1}{16}$

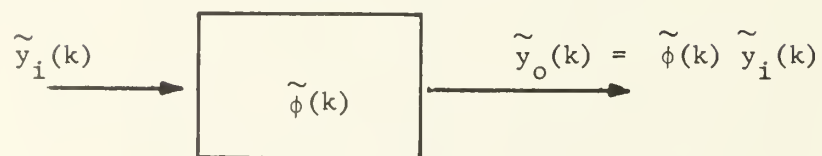


## LIST OF FIGURES

- Figure 1 - Filter Responses in the physical and transform domains.
- Figure 2 - Example of distortion by a filter.
- Figure 3 - Amplitude coefficients for the continuous filter.
- Figure 4 - (a) Phase and (b) group velocities for the continuous filter.
- Figure 5 - The discrete sampling schemes.
- Figure 6 - (a)-(c) Amplitude coefficients, (d) phase and (e) group velocities for the second-order finite difference filters as compared to the continuous (DIFF) filter.
- Figure 7 - (a)-(c) Amplitude coefficients, (d) phase and (e) group velocities for the fourth-order finite difference filters as compared to the continuous (DIFF) filter.
- Figure 8 - (a)-(c) Amplitude coefficients, (d) phase and (e) group velocities for the (piecewise linear) finite element filters as compared to the continuous (DIFF) filter.
- Figure 9 - Height field at  $t = 80$  hours as predicted by the continuous and discrete filter models.



(a) Physical Domain



(b) Transform Domain

Figure 1

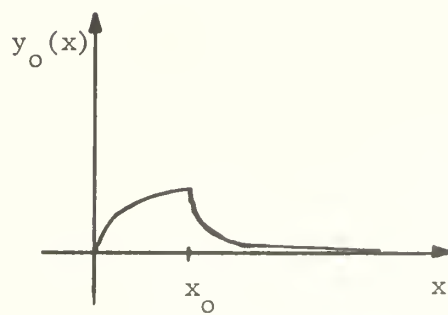
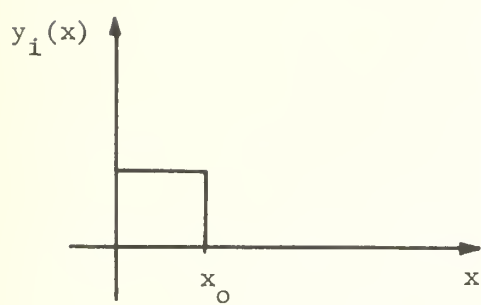
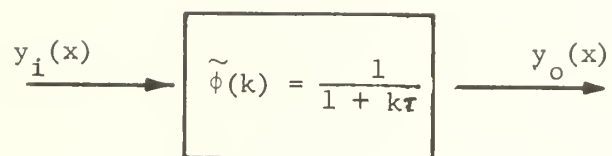


Figure 2

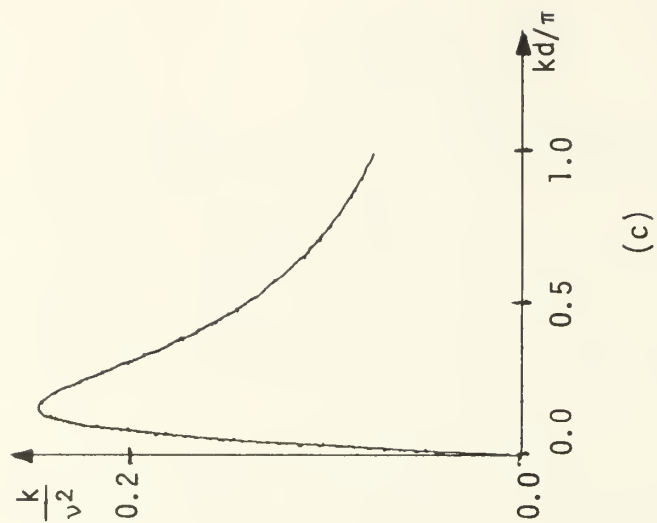
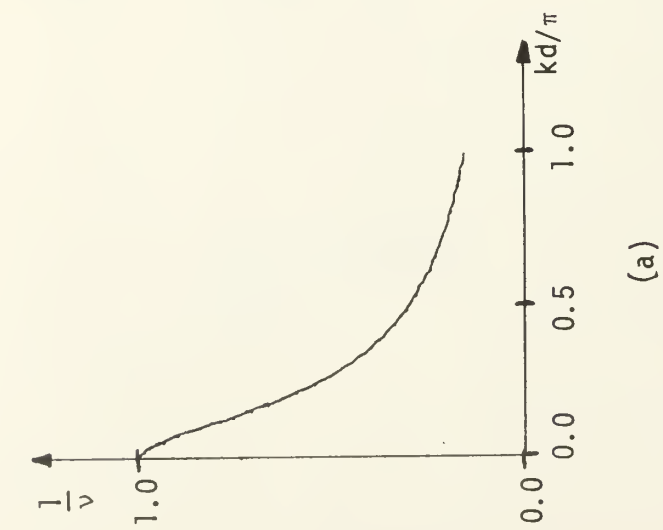
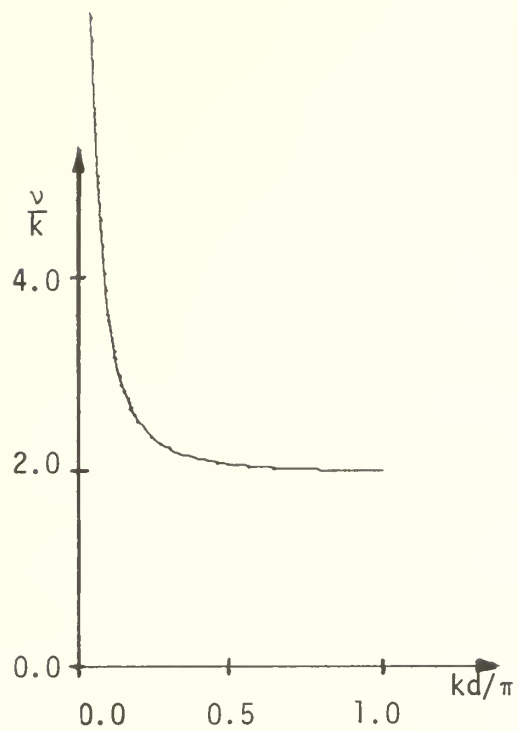
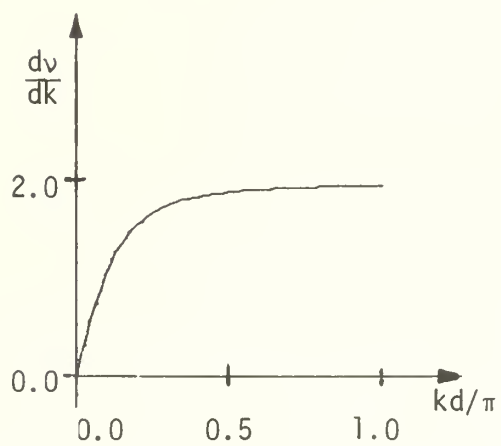


FIGURE 3

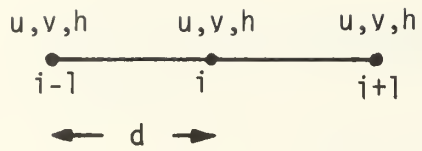


(a)

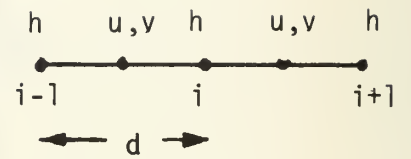


(b)

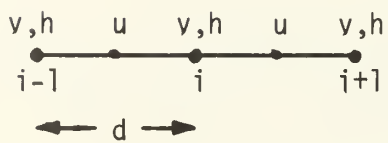
FIGURE 4



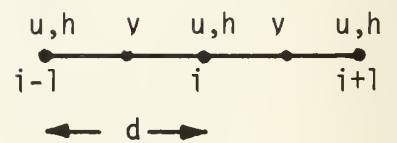
SCHEME A



SCHEME B



SCHEME C



SCHEME D

FIGURE 5

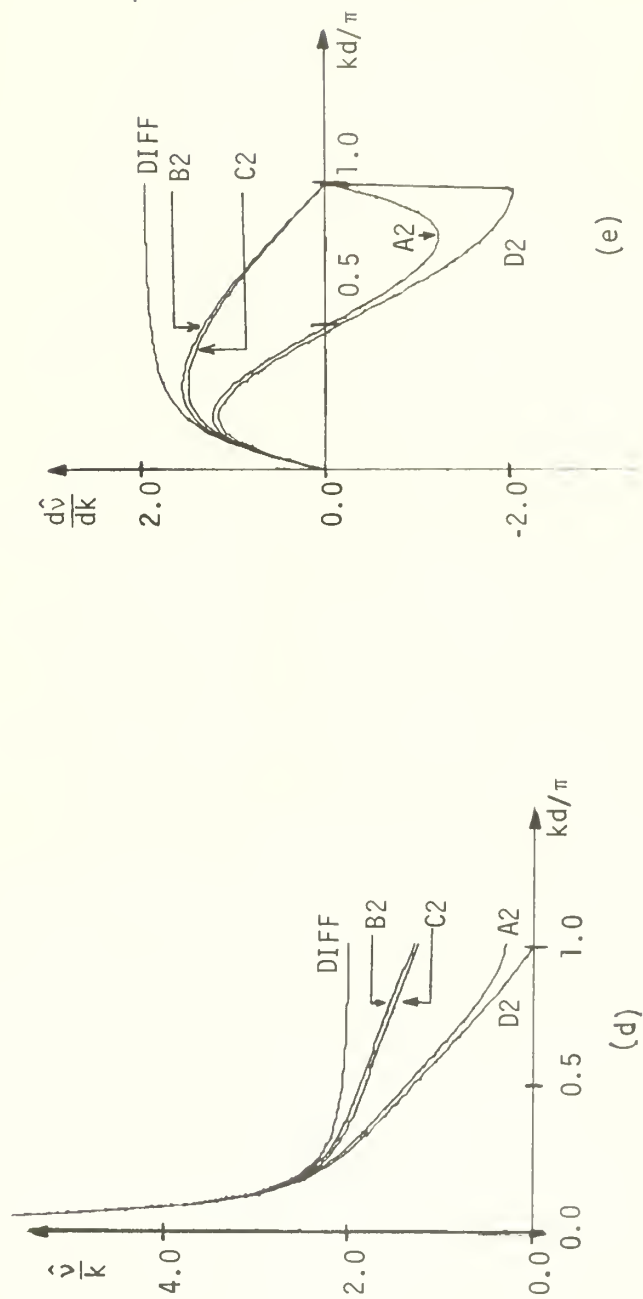
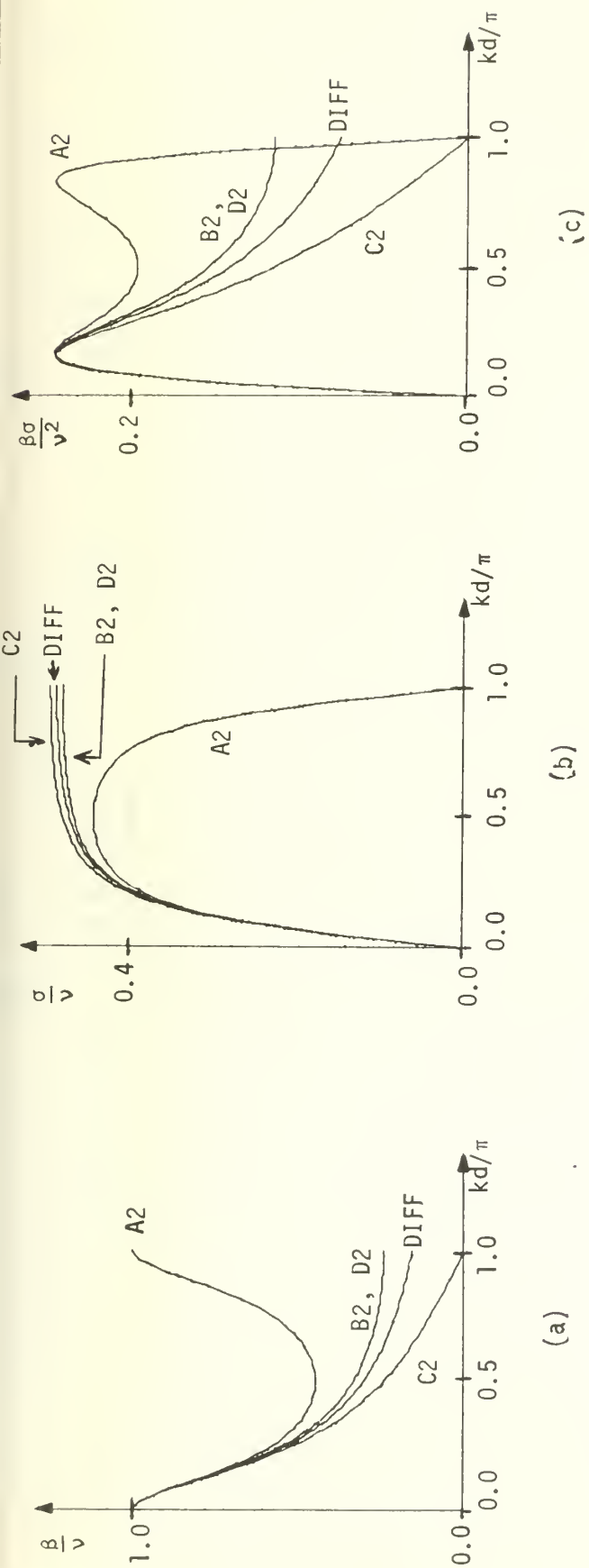
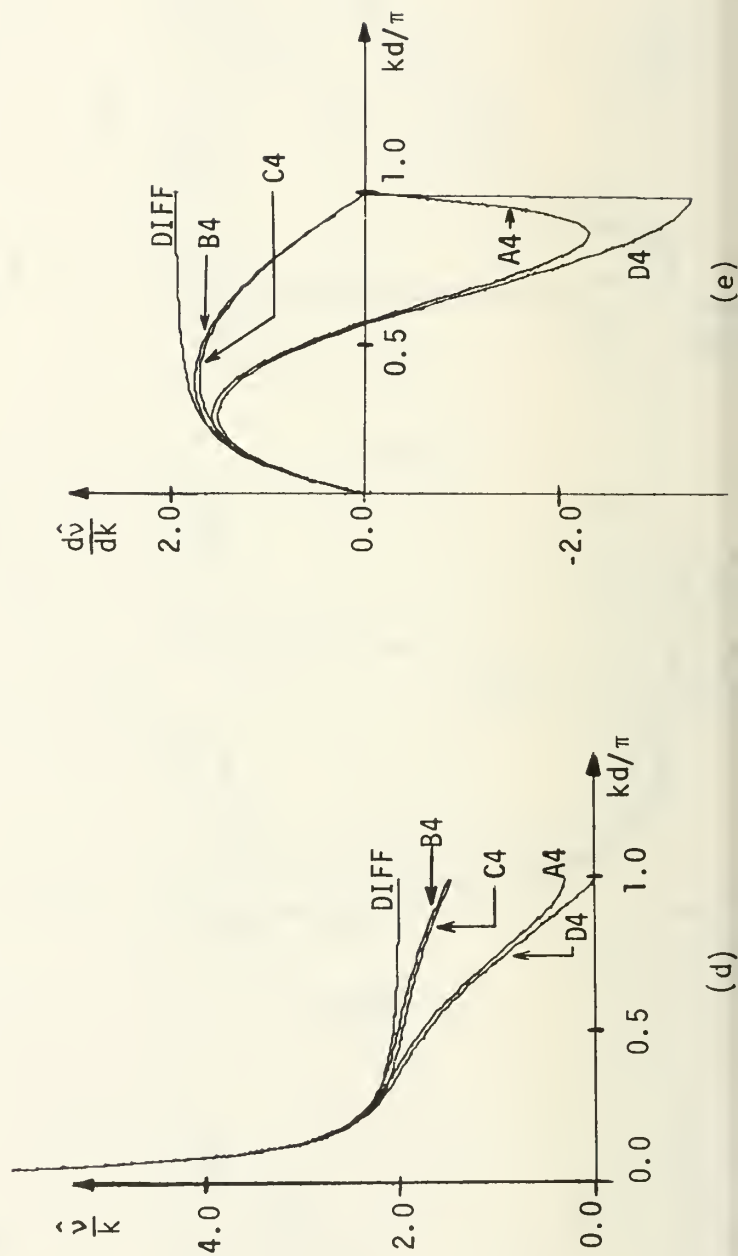
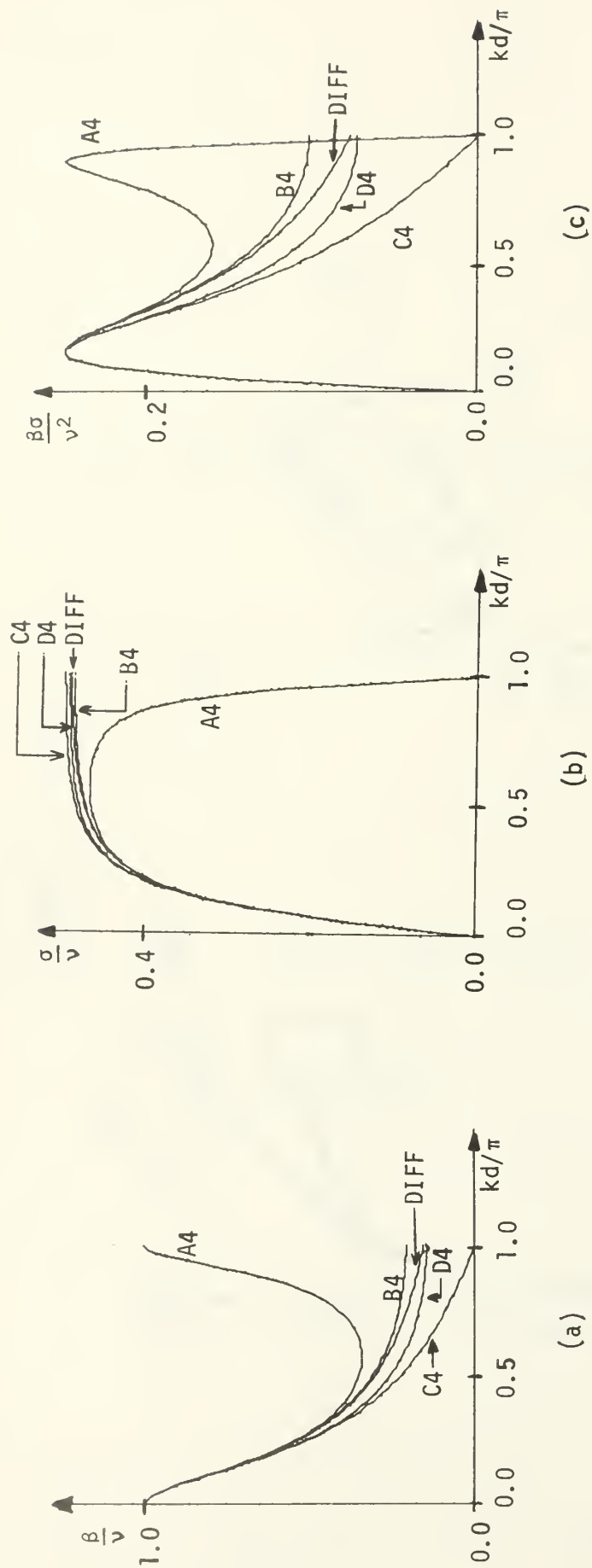


FIGURE 6





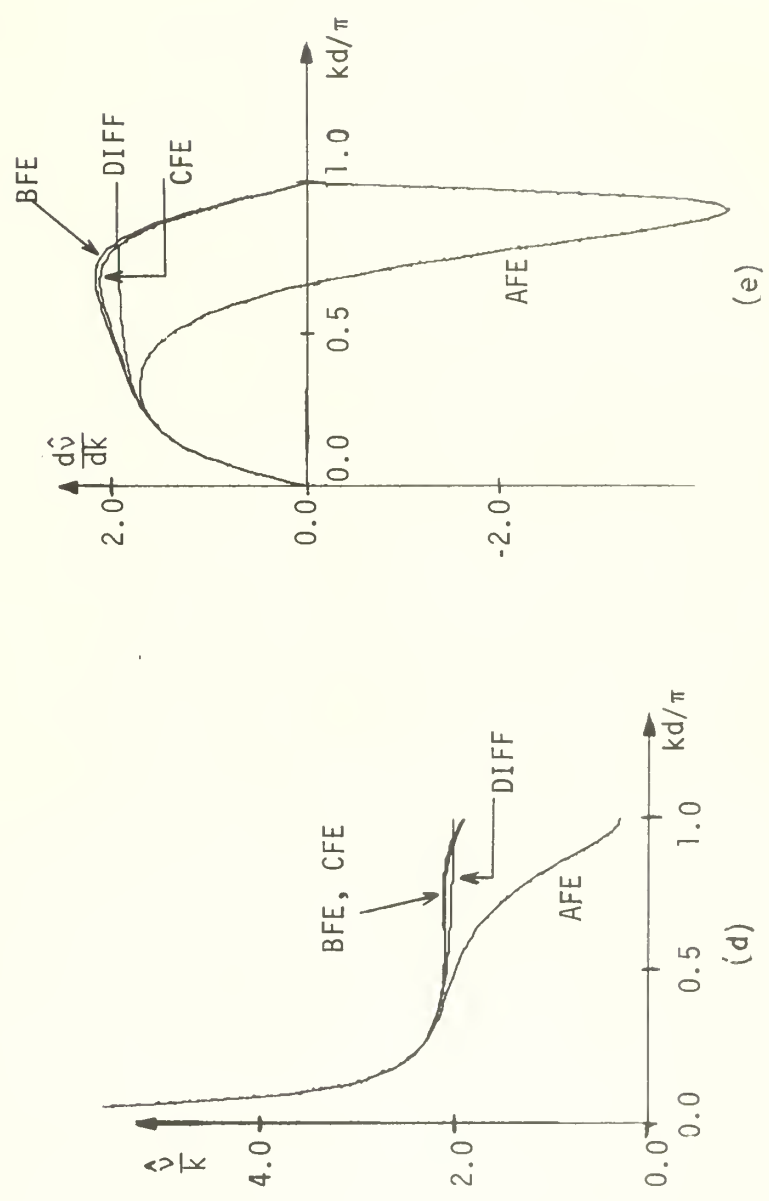
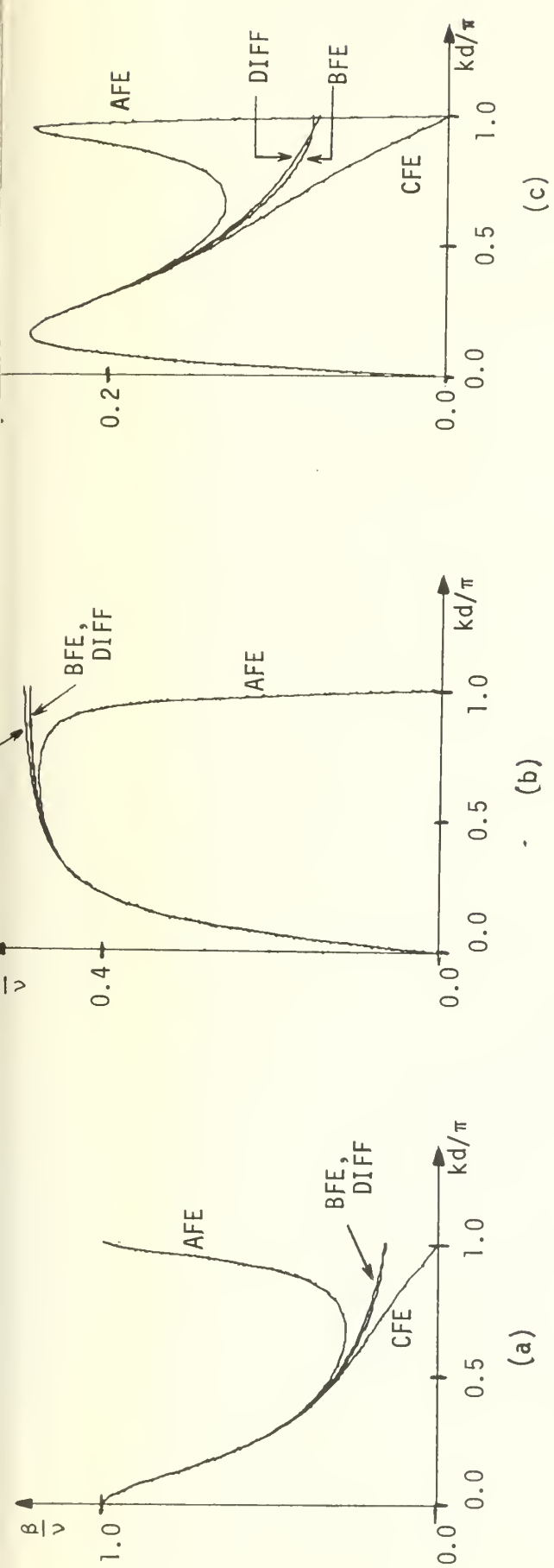


FIGURE 8

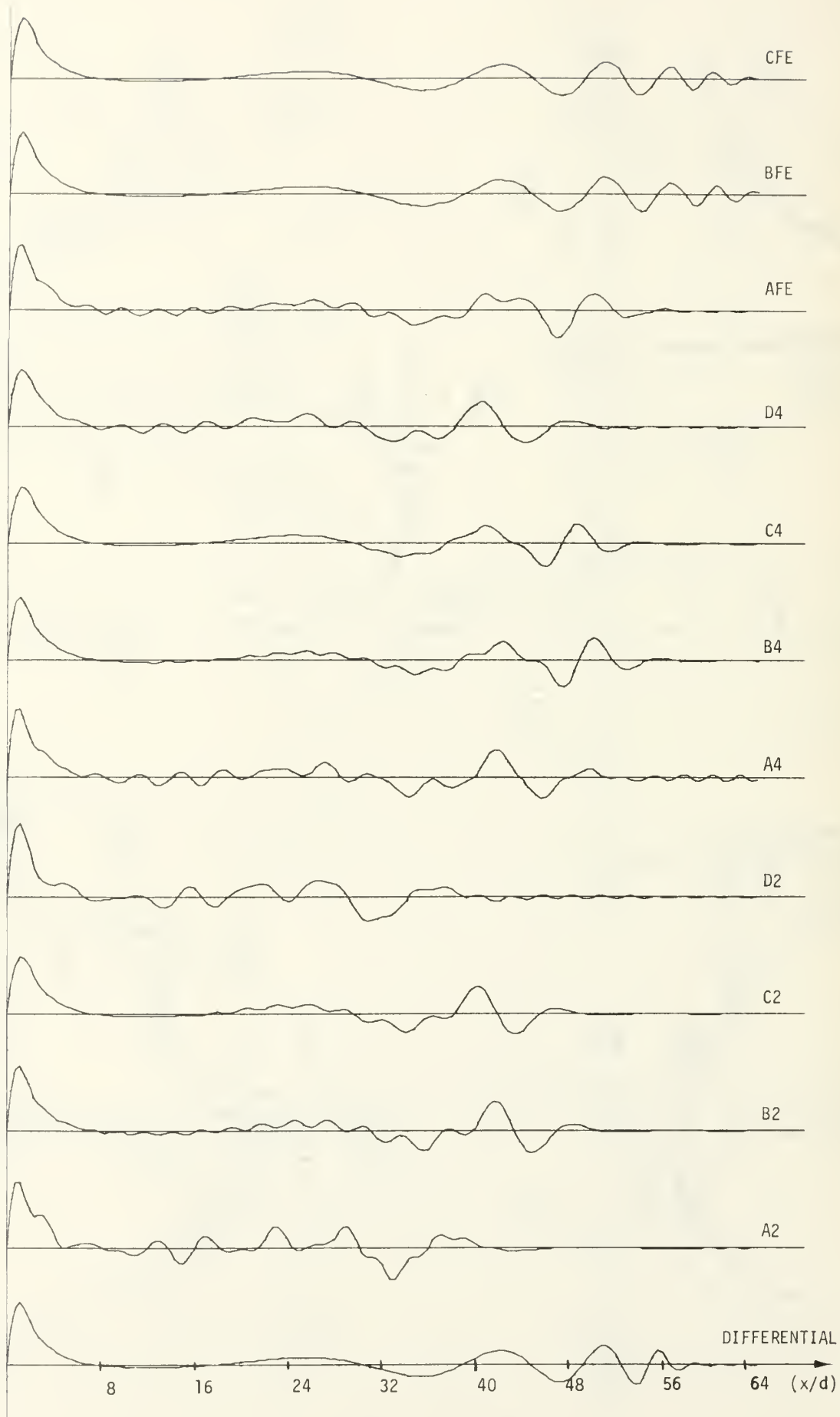


FIGURE 9

# DISTRIBUTION LIST

	No. Copies
1. Defense Documentation Center Cameron Station Alexandria, VA 22314	2
2. Dudley Knox Library, Code 0212 Naval Postgraduate School Monterey, CA 93940	2
3. Dr. A. L. Schoenstadt, Code 53Zh Department of Mathematics Naval Postgraduate School Monterey, CA 93940	26
4. Dr. R. T. Williams, Code 51Wu Department of Meteorology Naval Postgraduate School Monterey, CA 93940	1
5. Commanding Officer Naval Weather Service Command Headquarters 3101 Building 200 Washington Navy Yard Washington, D. C. 20374	1
6. Officer in Charge Environmental Prediction Research Facility Naval Postgraduate School Monterey, CA 93940	10
7. Dean of Research, Code 012 Naval Postgraduate School Monterey, CA 93940	2
8. Commanding Officer Fleet Numerical Weather Central Naval Postgraduate School Monterey, CA 93940	2
9. Naval Oceanographic Office Library (Code 3330) Washington, D. C. 20373	1
10. AFCL - Research Library L. G. Hanscom Field ATTN: Nancy Davis/Stop 29 Bedford, MA 01730	1

11. Commander, Air Weather Service 1  
Military Airlift Command  
United States Air Force  
Scott Air Force Base, IL 62226
12. Dr. R. Alexander 1  
The RAND Corporation  
1700 Main Street  
Santa Monica, CA 90406
13. Dr. A. Arakawa 1  
Department of Meteorology  
University of California  
Los Angeles, CA 90024
14. Dr. David A. Archer 1  
Douglas DuPont Rachford, Inc.  
6150 Chevy Chase  
Houston, TX 77027
15. Atmospheric Sciences Library 1  
National Oceanic and Atmospheric Administration  
Silver Spring, MD 20910
16. Mr. E. Barker 1  
Environmental Prediction Research Facility  
Naval Postgraduate School  
Monterey, CA 93940
17. Dr. Tom Beer 1  
Western Australian Institute of Technology  
Hayman Road South Bentley  
Western Australia 6102
18. Dr. W. Blumen 1  
Department of Astro-Geophysics  
University of Colorado  
Boulder, CO 80302
19. Dr. F. P. Bretherton 1  
National Center for Atmospheric Research  
P. O. Box 3000  
Boulder, CO 80303
20. Professor Dr. Jurgen Brickmann 1  
Universitat Konstanz  
Fachbereich Chemie  
775 Konstanz  
Postfach 77 33
21. Dr. John Brown 1  
National Meteorology  
Naval Postgraduate School  
Monterey, CA 93940

22. Dr. C. P. Chang, Code 51Cj 1  
 Department of Meteorology  
 Naval Postgraduate School  
 Monterey, CA 93940
  
23. Professor J. G. Charney 1  
 54-1424  
 Massachusetts Institute of Technology  
 Cambridge, MA 02139
  
24. Dr. Tom Delmer 1  
 Science Applications, Inc.  
 P. O. Box 2351  
 La Jolla, CA 92037
  
25. Dr. D. Dietrick 1  
 JAYCOR  
 205 S. Whiting St., Suite 409  
 Alexandria, VA 22304
  
26. Dr. Hugh W. Ellsaesser 1  
 Lawrence Livermore Laboratory  
 P. O. Box 808  
 Livermore, CA 94550
  
27. Dr. R. L. Elsberry, Code 63Es 1  
 Department of Meteorology  
 Naval Postgraduate School  
 Monterey, CA 93940
  
28. Dr. Frank D. Faulkner, Code 53Fa 1  
 Department of Mathematics  
 Naval Postgraduate School  
 Monterey, CA 93940
  
29. Dr. Richard Franke, Code 53Fe 1  
 Department of Mathematics  
 Naval Postgraduate School  
 Monterey, CA 93940
  
30. Dr. J. A. Galt 1  
 NOAA - Pac. Mar. Envir. Lab.  
 University of Washington WB-10  
 Seattle, WA 98105
  
31. Dr. W. L. Gates 1  
 The RAND Corporation  
 1700 Main Street  
 Santa Monica, CA 90406
  
32. Ian Gladwell 1  
 Department of Mathematics  
 University of Manchester  
 Manchester M13 9PL  
 ENGLAND

33. Dr. G. J. Haltiner, Code 63Ha 1  
Chairman, Department of Meteorology  
Naval Postgraduate School  
Monterey, CA 93940
  
34. Dr. R. L. Haney, Code 63Hy 1  
Department of Meteorology  
Naval Postgraduate School  
Monterey, CA 93940
  
35. LT Donald Hinsman 1  
Fleet Numerical Weather Central  
Naval Postgraduate School  
Monterey, CA 93940
  
36. D. Narayana Holla 1  
Department of Maths  
Indian Institute of Technology, Bombay  
P. O., I. I. T. Bombay  
Powai, Bombay 400 076 INDIA
  
37. Dr. J. Holton 1  
Department of Atmospheric Sciences  
University of Washington  
Seattle, WA 98105
  
38. Dr. B. J. Hoskins 1  
Department of Geophysics  
University of Reading  
Reading  
United Kingdom
  
39. W. Horsthemke 1  
Universite Libre De Bruxelles-Faculte Des Sci.  
Service De Chimie Physique II-Prof. Prigogine  
C.P. 231 Campus Plaine-Bd. du Triomphe  
1050 Bruxelles - BELGIQUE
  
40. Dr. D. Houghton 1  
Department of Meteorology  
University of Wisconsin  
Madison, WI 53706
  
41. Dennis C. Jespersen 1  
Mathematics Research Center  
University of Wisconsin-Madison  
610 Walnut Street  
Madison, WI 53706
  
42. Dr. A. Kasahara 1  
National Center for Atmospheric Research  
P. O. Box 3000  
Boulder, CO 80303

43. E. J. Kansh 1  
U.S. Department of the Interior  
Bureau of Mines  
Pittsburgh, PA 15213
  
44. Dr. L. D. Kovach, Code 53Kv 1  
Department of Mathematics  
Naval Postgraduate School  
Monterey, CA 93940
  
45. Ing. J. Kumicak 1  
Institute of Radioecology  
and Applied Nuclear Techniques  
Komenskeho 9  
P. O. Box A-41  
040 61 Kosice  
CSSR CZECHOSLOVAKIA
  
46. Dr. C. E. Leith 1  
National Center for Atmospheric Research  
P. O. Box 3000  
Boulder, CO 80303
  
47. Dr. J. M. Lewis 1  
Laboratory for Atmospheric Research  
University of Illinois  
Urbana, IL 61801
  
48. Mei-Kao Liu 1  
System Applications, Inc.  
950 Northgate Drive  
San Rafael, CA 94903
  
49. Dr. E. N. Lorenz 1  
Department of Meteorology  
Massachusetts Institute of Technology  
Cambridge, MA 02139
  
50. F. Ludwikow 1  
Medical School Department of Biophysics  
ul. Chalubinskiego 10  
50-368 Wroclaw  
POLAND
  
51. Leon Lupidus 1  
Princeton University  
School of Engineering/Applied Science  
Department of Chemical Engineering  
The Engineering Quadrangle  
Princeton, NJ 08540

52. Dr. R. Madala 1  
Code 7750  
Naval Research Laboratory  
Washington, D. C. 20390
53. Dr. J. D. Mahlman 1  
Geophysical Fluid Dynamics Laboratory  
Princeton University  
Princeton, NJ 08540
54. Dr. Alsan Meric 1  
Applied Mathematics Division  
Marmara Research Institute  
P.K. 141, Kadikoy - Istanbul  
TURKEY
55. Dr. G. Morris, Code 53Mj 1  
Department of Mathematics  
Naval Postgraduate School  
Monterey, CA 93940
56. Professor J. LL. Morris 1  
Computer Science  
University of Waterloo  
Waterloo, Ontario  
CANADA
57. Meteorology Library (Code 63) 1  
Naval Postgraduate School  
Monterey, CA 93940
58. LT W. F. Mihok 1  
Fleet Numerical Weather Central  
Naval Postgraduate School  
Monterey, CA 93940
59. National Center for Atmospheric Research 1  
Box 1470  
Boulder, CO 80302
60. Director, Naval Research Laboratory 1  
ATTN: Technical Services Information Center  
Washington, D. C. 20390
61. Department of Oceanography, Code 68 1  
Naval Postgraduate School  
Monterey, CA 93940
62. Office of Naval Research 1  
Department of the Navy  
Washington, D. C. 20360



63. Dr. T. Ogura 1  
Laboratory for Atmospheric Research  
University of Illinois  
Urbana, IL 61801
  
64. Darrell W. Pepper 1  
Environ. Transport Div.  
E.I. du Pont de Nemours & Co., Inc.  
Savannah River Laboratory  
Aiken, SC 29801
  
65. Professor N. A. Phillips 1  
National Meteorological Center/NOAA  
World Weather Building  
Washington, D. C. 20233
  
66. Dr. S. Piacsek 1  
Code 7750  
Naval Research Laboratory  
Washington, D. C. 20390
  
67. A. P. Raiche 1  
Minerals Research Laboratories, C.S.I.R.O.  
P. O. Box 136  
North Ryde, N.S.W. 2113  
AUSTRALIA
  
68. Dr. T. Rosmond 1  
Environmental Prediction Research Facility  
Naval Postgraduate School  
Monterey, CA 93940
  
69. Dr. Y. Sasaki 1  
Department of Meteorology  
University of Oklahoma  
Norman, OK 73069
  
70. Dr. Fred Shuman, Director 1  
National Meteorological Center  
World Weather Building  
Washington, D. C. 20233
  
71. Dr. R. Somerville 1  
NCAR  
P. O. Box 3000  
Boulder, CO 80303
  
72. Dr. Mevlüt Teymur 1  
Applied Mathematics Division  
Marmara Research Institute  
P.K. 141, Kadikoy - Istanbul  
TURKEY

73. Dr. D. Williamson 1  
National Center for Atmospheric Research  
P. O. Box 3000  
Boulder, CO 80303
74. Professor Carroll O. Wilde 1  
Chairman, Department of Mathematics  
Naval Postgraduate School  
Monterey, CA 93940
75. Dr. F. J. Winninghoff 1  
1085 Steeles Avenue, #503  
Willowdale (Toronto)  
Ontario M2R2T1 CANADA
76. Dr. M. G. Wurtele 1  
Department of Meteorology  
University of California  
Los Angeles, CA 90024
77. Dr. J. Young 1  
Department of Meteorology  
University of Wisconsin  
Madison, WI 53706

U186021

DUDELEY WOOD LIBRARY - RESEARCH REPORTS

5 6853 01057937 8

~~41860~~